Evidence of Early Supershear Transition in the Feb 6th 2023 M_w 7.8 Kahramanmaraş Turkey Earthquake From Near-Field Records

Ares Rosakis^a, Mohamed Abdelmeguid^a, Ahmed Elbanna^{b,c}

^aGraduate Aerospace Laboratories, California Institute of Technology, Pasadena, CA, ^bDepartment of Civil and Environmental Engineering, University of Illinois at Urbana Champaign, Urbana, IL, ^cBeckman Institute of Advanced Science and Technology, University of Illinois at Urbana Champaign, Urbana, IL,

4 Abstract

3

The $M_w7.8$ Kahramanmaraş Earthquake was larger and more destructive than what had been expected for the tectonic setting in Southeastern Turkey. By using near-field records we provide evidence for early supershear transition on the splay fault that hosted the nucleation and early propagation of the first rupture that eventually transitioned into the East Anatolian fault. We also find, for the first time ever, field observational evidence showing the mechanism of sub-Rayleigh to supershear transition. We estimate the instantaneous supershear rupture propagation speed to be $\sim 1.55C_s$ and the sub-Rayleigh to supershear transition length to be around ~ 19.45 km, very close to the location of one of the stations, closest to the epicenter. This early supershear transition might have facilitated the continued propagation and triggering of slip on the nearby East Anatolian Fault leading to amplification of the hazard. The complex dynamics of the Kahramanmaraş earthquake warrants further studies.

5 Introduction

⁶ On February 6th 2023, a M_w 7.8 earthquake shook the southeastern parts of Turkey and ⁷ northern Syria. Preliminary back projection models based on teleseismic data as well as mul-⁸ tiple seismic inversions suggest that rupture initiated at 1:17:355 coordinated universal time ⁹ (UTC) on a splay branch fault in the near proximity of the East Anatolian fault [1]. The precise

location of the hypocenter is currently uncertain. The preliminary hypocenter location was esti-10 mated by AFAD to be 37.288°N 37.042°E [2] with a depth of approximately 8 km. It was also 11 estimated by the USGS to be 37.166°N 37.042°E \pm 6.3 km (indicated by the red star marker in 12 Figure 1) with a depth of approximately 18 ± 3 km [1]. The rupture then propagated north east 13 subsequently transferring to the East Anatolian fault and starting a sequence of seismic events. 14 Furthermore, subsequent preliminary geodetic inversions confirmed the multi-segment nature 15 of the M_w 7.8 rupture. The sequence of events resulted in catastrophic levels of destruction 16 with substantial humanitarian and financial losses. Based on historical records, the magnitude 17 of the event and the total rupture length were both much larger than expected for such a tectonic 18 setting in southern Turkey [3]. This together with the intensity of the measured ground shaking 19 motivated us to investigate the nature of rupture initiation, propagation, as well as the possibility 20 of early supershear transition. 21

Figure 1 illustrates the estimated location of the hypocenter, the approximate strike of the 22 splay fault which is inferred to be around N22°E based on the aftershock sequence, and the 23 sense of motion (left lateral for both the splay fault, and the east Anatolian fault). To the best 24 of our knowledge, three stations exist very close to the splay fault as highlighted by the green 25 diamonds in Figure 1. Two of these stations: TK:NAR and KO:KHMN are located at 37.3919°N 26 37.1574°E [2, 4], and herein are referred to as the twin stations because they are at the same 27 geographical location. Another station TK:4615 is located closer to the epicenter at 37.386°N 28 37.138°E [2]. The insert in Figure 1 is a schematic of the positions of the stations, showing 29 the distances x_1, x_2 relative to the epicenter and the distances L_1, L_2 relative to the hypocenter 30 which is located at a depth d. These three stations provide a rare and detailed insight into the 31 near-field characteristics of the rupture on the splay fault and indeed close examination of these 32 records have revealed unique observations that we describe below. 33

34 Clear signature of supershear in the twin stations records

Figure 2a shows the time histories of the particle velocities along the fault parallel, the fault 35 normal, and the vertical directions from the twin stations (TK:NAR solid black line, KO:KHMN 36 solid red line). These are obtained from the instrument corrected ground motions. The raw 37 NS, EW and vertical acceleration records are obtained from (AFAD) and (KOERI) respec-38 tively (Retrieved 02/09 5:18 PST) [2, 4]. We computed the velocities for TK:NAR by numer-39 ically integrating the available acceleration records from AFAD [2]. The velocity response for 40 KO:KHMN was processed using the Obspy software [5]. We then resolved the computed NS 41 and EW ground velocity signals parallel and perpendicular to the splay fault shown in Figure 42 1. To the best of our knowledge, these records correspond to two different instruments and as 43 a result the good agreement between the records provides a degree of confidence in the quality 44 of the data to be used in the present study. Here, the first vertical dashed line indicates the first 45 arrival of P-waves from the hypocenter based on the rupture initiation at the USGS provided 46 time 1:17:355 coordinated universal time (UTC) [1]. 47

The velocity waveforms for the twin stations reveal unique characteristics. We first observe 48 that the FP component is clearly more dominant than the FN component. This is atypical of sub-49 Rayleigh strike-slip earthquake ruptures which feature more dominant fault normal versus fault 50 parallel velocity components. However, a dominant fault parallel component is a characteristic 51 feature of supershear ruptures [6, 7] in which the rupture speed exceeds the shear wave speed of 52 crustal rock C_s . Such a behavior has been observed both in the laboratory [8, 9, 10] and the field 53 [9, 11, 12, 13], and has been also predicted by the theory [8, 11, 14]. This provides evidence for 54 supershear rupture propagation towards the twin stations. 55

⁵⁶ We observe intense ground shaking associated with the arrival of the supershear Mach cone ⁵⁷ at the station and we identify this arrival by the red dashed line. Through measuring the change ⁵⁸ in ground motion associated with the supershear Mach front, we observe that the ratio of the fault parallel $\delta \dot{u}_{FP}^s$ to the fault normal component $\delta \dot{u}_{FN}^s$ is approximately ~ 1.2. As discussed by Mello et al. 2016, these changes correspond to the shear part of the velocity signal, and are due to the arrival of the shear Mach lines [8]. The ratio of the changes in the particle velocities has been theoretically shown by Mello et al. 2016 to depend uniquely on the ratio of the rupture speed and the shear wave speed as follows $\delta \dot{u}_{FP}^s / \delta \dot{u}_{FN}^s = \sqrt{(V_r/C_s)^2 - 1}$. This relationship is also shown schematically in Figure 2b. Accordingly, and as indicated in the figure, for a ratio of 1.2, the corresponding supershear rupture speed is ~ $1.55C_s$.

Furthermore, in Figure 2a, the black dashed line indicates the eventual arrival of the trailing Rayleigh signature which represents the remnant of the initially sub-Rayleigh rupture before it transitioned to supershear. Figure 2c is a top view detailing the location of the three stations relative to the epicenter, highlighting the transition length L_T after which the rupture speed V_r exceeds the shear wave speed C_s . It also shows the shear Mach cone interaction with the stations.

Based on the geometry of Figure 2c, and assuming that the rupture tip initially propagates at $V_r = C_r$ prior to transition between (0,0) and $(0, L_T)$ and then transition to $V_r = 1.55C_s$ till it arrives at the twin stations at x_2 , we can estimate a transition length L_T by further assuming that the stations are located on the fault [9, 15].

$$L_T = C_R \frac{x_2 - t_s V_r}{C_R - V_r} \tag{1}$$

⁷⁶ Where, t_s is the arrival time of the shear Mach cone to the station which can be obtained ⁷⁷ from Figure 2a (red dashed line), and V_r is the supershear rupture speed $1.55C_s$. In the above ⁷⁸ relationship, x_2 is furnished as $\sqrt{L_2^2 - d^2}$ as shown in the insert of Figure 1, where L_2 is the ⁷⁹ distance of the twin stations from the hypocenter at depth d. L_2 is estimated based on the ⁸⁰ P-arrival time (first disturbance) from the hypocenter location to the station, and the assumed ⁸¹ dilatational wave speed C_p as we will describe shortly.

82 Evidence of sub–Rayleigh to supershear transition in the TK:4165 station record

Similar to Figure 2a, Figure 3a shows the time histories of the particle velocities along 83 the fault parallel, the fault normal, and the vertical directions obtained from station TK:4165 84 (AFAD) [2]. However, this record is qualitatively different from the record shown in Figure 2a. 85 Indeed, we observe here that the fault normal velocity component is larger than the fault par-86 allel component, which is characteristic of a primarily sub-Rayleigh rupture. However, careful 87 examination of the fault parallel record indicates the presence of a small but well defined pulse 88 ahead of the Rayleigh signature as indicated in the top panel of Figure 3a (shaded region). We 89 believe that this feature is a supershear pulse, which has just been formed ahead of the primary 90 rupture which is still propagating at the Rayleigh wave speed. Accordingly, we hypothesize 91 that station TK:4165 is located very close to the point where the rupture transitioned from sub-92 Rayleigh to supershear. It should be noted that the probability of capturing the early stages of 93 Rayleigh to supershear rupture transition is very low, and has never been observed before in a 94 near fault field record. However, this transition has been reported experimentally in laboratory 95 earthquakes performed by Rosakis et al 2004 [16] and Mello et al 2016 [8] (We refer the reader 96 to Figure 14 in [8] for illustration). Specifically, Mello et al 2016 captured this transition by 97 comparing dynamic, full field photoelastic images of the initial stages of the formation of the 98 supershear pulse with near fault particle velocity records measured at a location close to the 99 transitioning rupture and by further correlating the two measurement techniques. The velocity 100 records were obtained experimentally by a pair of laser velocimeters recording the fault parallel 101 and fault normal components [8]. 102

To investigate the validity of this hypothesis, related to supershear transition and the location of TK:4615, we present a preliminary analysis by comparing the location of the station x_1 to our independent estimate of L_T obtained from the twin stations record shown in Figure 2a. In order to do this, we assume $C_s = 3320$ m/s, and $C_p = 5780$ m/s which correspond to a Poisson's ratio

of 0.25, and are in good agreement with velocity models for the southern Turkey region [3]. It 107 follows then that $C_R = 3050$ m/s and $V_r = 5146$ m/s. Based on the P-arrival time at the twin 108 stations and using the above C_p leads to $L_2 = 23.7$ km. We note that for a hypocenter depth of 109 d = 10.9 km, equation (1) yields a transition length $L_T = 19.45$ km. We then use the P-wave 110 arrival time at station TK:4165 to identify its distance from the hypocenter $L_1 = 22.3$ km. Using 111 the Pythagorean theorem, we compute the epicentral distance of station TK:4165 as $x_1 = 19.45$ 112 km. For this particular choice of depth d, we observe that the location of the station TK:4165 113 coincides with the location of the sub-Rayleigh to supershear transition, which is consistent 114 with our hypothesis. This estimate of depth of 10.9 km is within the range predicted by the 115 different agencies (AFAD and USGS) [1, 2]. Furthermore, computing the distance between the 116 twin stations and TK:4165 yields $\delta x = x_2 - x_1 = 1.6$ km along the fault strike direction. Since 117 the total distance between the twin stations and TK:4165 is ~ 2 km, based on their respective 118 coordinates, this computed difference in their epicentral distances is a plausible estimate. 119

120 Discussion

Our analysis of three rare near-field (~ 1 km from the fault) velocity records of the $M_w 7.8$ 121 Kahramanmaraş earthquake suggests the rupture that propagated on the splay fault had tran-122 sitioned from sub-Rayleigh to supershear speed ($V_r \sim 1.55 C_s$) at an epicentral distance of 123 approximately 19.45 km. The records obtained from the twin stations showing perfect agree-124 ment with one another provides confidence in the quality of the data to be used in the present 125 study. In addition, a station located in such near proximity to the transition point is a unique 126 occurrence, that to our best knowledge has never been reported before in the literature. Those 127 rare near-field records captured, for the first time, the in-situ transition mechanism from sub-128 Rayleigh to supershear propagation and provided a detailed window into the structure of the 129 near-fault particle motions in both the fault parallel and fault normal directions. It is unprece-130

dented to have multiple near-field stations capturing the field dynamics of supershear rupture 131 transition and propagation. This makes these records particularly important and emphasizes the 132 value of having high-quality near-field data, as such data carries significant local information 133 about the rupture physics which may be lost in the far-field measurements [17]. Furthermore, 134 since Mach fronts attenuate only weakly with distance, this early supershear transition on the 135 splay fault may have enabled strong dynamic stress transfer to the nearby East Anatolian Fault 136 and contributed to the continued rupture propagation and triggered slip in both the North East 137 and South West directions as in previous earthquakes [18]. Indeed, prior studies have suggested 138 that supershear ruptures are more effective in jumping across fault stepovers [19] and activa-139 tion of nearby faults [20, 21, 22, 23]. The early supershear transition on the splay fault may 140 have been favored by the regional stress state. Seismological studies suggest that the splay fault 141 exists in a N16.4°E compression regime (σ_1) and it is under the N80.8°W extension regime 142 (σ_3) [24]. The estimated strike of the splay fault N22°E thus makes it close to being perpendic-143 ular to the direction of the minimum principal stress which reduces the overall normal stress on 144 the fault. This may significantly reduce the fault strength parameter S (e.g. S<1) [25, 16] and 145 favors transition to supershear rupture over shorter distances. Other mechanisms that may have 146 favored a rapid supershear transition include on-fault stress or strength heterogeneities [26, 27] 147 or off-fault material complexities [28, 29]. The extended propagation of the rupture in the NNE 148 direction may also suggest the existence of a velocity contrast across the fault surface and a 149 bimaterial effect[30, 31, 32]. Overall, we hope that further studies of the regional stress field 150 and the structure of the ground motion records will reveal more details about the nature of this 151 complex multi-segment rupture that led to such a large-scale human tragedy. Future detailed 152 numerical simulations and analog experimental investigations are also needed to better con-153 strain the dynamics of complex fault zones, like the East Anatolian Fault Zone, beyond what is 154 available from historical records and regional scaling relations. This will help reduce the impact 155

¹⁵⁶ of future hazards and better inform preparedness efforts.

157 Acknowledgement

A.J.R. acknowledges support by the Caltech/MCE Big Ideas Fund (BIF), as well as the 158 Caltech Terrestrial Hazard Observation and Reporting Center (THOR). He would also like to 159 acknowledge the support of NSF (Grant EAR-1651235 and EAR-1651235). A.E. acknowledge 160 support from the Southern California Earthquake Center through a collaborative agreement 161 between NSF. Grant Number: EAR0529922 and USGS. Grant Number: 07HQAG0008 and the 162 National Science Foundation CAREER award No. 1753249 for modeling complex fault zone 163 structures. The ground motion data used in this study can be obtained from Turkish Disaster 164 and Emergency Managment Authority AFAD, US Geological Survey (USGS), and Kandilli 165 Observatory And Earthquake Research Institute. We would like to thank the Turkish Disaster 166 and Emergency Management Presidency (AFAD) for setting up dense near-fault observatories, 167 and for immediately publishing a huge number of openly accessible accelerometers during these 168 trying times for Turkey. 169

170 **References**

- [1] US Geological Survey, M 7.8 27 km E of Nurdağı, Turkey (2).
- URL https://earthquake.usgs.gov/earthquakes/eventpage/ us6000jllz/executive
- 174 [2] Disaster, E. M. Authority, Turkish National Strong Motion Network (1973). doi:
- 175 https://doi.org/10.7914/SN/TK.
- 176 URL https://tadas.afad.gov.tr
- [3] D. Acarel, M. D. Cambaz, F. Turhan, A. K. Mutlu, R. Polat, Seismotectonics of Malatya

178	Fault, Eastern Turkey, Open Geosciences 11 (1) (2019) 1098–1111.	doi:10.1515/
179	geo-2019-0085.	

[4] Kandilli Observatory And Earthquake Research Institute Boğaziçi University, Kandilli
 Observatory And Earthquake Research Institute (KOERI) (1971). doi:10.7914/SN/
 KO.

183 URL https://www.fdsn.org/networks/detail/KO/

- [5] M. Beyreuther, R. Barsch, L. Krischer, T. Megies, Y. Behr, J. Wassermann, ObsPy: A
 Python Toolbox for Seismology, Seismological Research Letters 81 (3) (2010) 530–533.
 doi:10.1785/gssrl.81.3.530.
- 187 URL https://pubs.geoscienceworld.org/srl/article/81/3/ 188 530-533/143693
- [6] A. J. Rosakis, O. Samudrala, D. Coker, Cracks Faster than the Shear Wave Speed, Science
 284 (5418) (1999) 1337–1340. doi:10.1126/science.284.5418.1337.
- 191 URL https://www.sciencemag.org/lookup/doi/10.1126/science.
 192 284.5418.1337
- [7] M. Bouchon, H. Karabulut, M. P. Bouin, J. Schmittbuhl, M. Vallée, R. Archuleta, S. Das,
 F. Renard, D. Marsan, Faulting characteristics of supershear earthquakes, Tectonophysics
 493 (3-4) (2010) 244–253. doi:10.1016/j.tecto.2010.06.011.
- [8] M. Mello, H. S. Bhat, A. J. Rosakis, Spatiotemporal properties of Sub-Rayleigh and
 supershear rupture velocity fields: Theory and experiments, Journal of the Mechanics and
 Physics of Solids 93 (2016) 153–181. doi:10.1016/j.jmps.2016.02.031.
- 199 URL http://dx.doi.org/10.1016/j.jmps.2016.02.031https:
- 200 //linkinghub.elsevier.com/retrieve/pii/S0022509616301363

- [9] M. Mello, H. S. Bhat, A. J. Rosakis, H. Kanamori, Reproducing the supershear portion of
 the 2002 Denali earthquake rupture in laboratory, Earth and Planetary Science Letters 387
- 203 (2014) 89-96. doi:10.1016/j.epsl.2013.11.030.
- 204 URL http://dx.doi.org/10.1016/j.epsl.2013.11.030
- [10] X. Lu, A. J. Rosakis, N. Lapusta, Rupture modes in laboratory earthquakes: Effect of
 fault prestress and nucleation conditions, Journal of Geophysical Research: Solid Earth
 115 (12) (2010) 1–25. doi:10.1029/2009JB006833.
- [11] E. M. Dunham, R. J. Archuleta, Evidence for a supershear transient during the 2002 Denali fault earthquake, Bulletin of the Seismological Society of America 94 (6 SUPPL. B)
 (2004) 256–268. doi:10.1785/0120040616.
- [12] M. Bouchon, M.-P. Bouin, H. Karabulut, M. N. Toksöz, M. Dietrich, A. J. Rosakis,
 How fast is rupture during an earthquake? New insights from the 1999 Turkey Earthquakes, Geophysical Research Letters 28 (14) (2001) 2723–2726. doi:10.1029/
 2001GL013112.
- ²¹⁵ URL http://doi.wiley.com/10.1029/2001GL013112

[13] H. Zeng, S. Wei, A. Rosakis, A Travel-Time Path Calibration Strategy for Back-Projection
 of Large Earthquakes and Its Application and Validation Through the Segmented Super Shear Rupture Imaging of the 2002 Mw 7.9 Denali Earthquake, Journal of Geophysical
 Research: Solid Earth 127 (6). doi:10.1029/2022JB024359.

- [14] E. M. Dunham, H. S. Bhat, Attenuation of radiated ground motion and stresses from three-
- dimensional supershear ruptures, Journal of Geophysical Research: Solid Earth 113 (B8)

222 (2008) 1-17. doi:10.1029/2007JB005182.

223 URL http://doi.wiley.com/10.1029/2007JB005182

- [15] V. Rubino, A. J. Rosakis, N. Lapusta, Spatiotemporal Properties of Sub-Rayleigh and 224 Supershear Ruptures Inferred From Full-Field Dynamic Imaging of Laboratory Exper-225 iments, Journal of Geophysical Research: Solid Earth 125 (2) (2020) 1-25. doi: 226 10.1029/2019JB018922. 227
- A. J. Rosakis, H. Kanamori, Laboratory Earthquakes: The Sub-[16] K. Xia, 228 Rayleigh-to-Supershear Rupture Transition, Science 303 (5665) (2004) 1859–1861. 229 doi:10.1126/science.1094022. 230
- https://www.sciencemaq.org/lookup/doi/10.1126/science. URL 231 1094022 232
- [17] Y. Ben-Zion, A Critical Data Gap in Earthquake Physics, Seismological Research Letters 233 90(5)(2019)1721-1722.doi:10.1785/0220190167. 234
- URL https://pubs.geoscienceworld.org/ssa/srl/article/ 235 572859/A-Critical-Data-Gap-in-Earthquake-Physics 236
- [18] S. Das, The Need to Study Speed, Science 317 (5840) (2007) 905–906. doi:10.1126/ 237
- science.1142143. 238

243

- URL https://www.science.org/doi/10.1126/science.1142143 239
- [19] R. A. Harris, S. M. Day, Dynamics of fault interaction: parallel strike-slip faults, Journal 240 of Geophysical Research 98 (B3) (1993) 4461-4472. doi:10.1029/92JB02272. 241
- [20] E. L. Templeton, A. Baudet, H. S. Bhat, R. Dmowska, J. R. Rice, A. J. Rosakis, C. E. 242
- Rousseau, Finite element simulations of dynamic shear rupture experiments and dynamic
- path selection along kinked and branched faults, Journal of Geophysical Research: Solid 244
- Earth 114 (8). doi:10.1029/2008JB006174. 245

246	[21]	C. E. Rousseau, A. J. Rosakis, Dynamic path selection along branched faults: Experiments
247		involving sub-Rayleigh and supershear ruptures, Journal of Geophysical Research: Solid
248		Earth 114 (8) (2009) 1-15. doi:10.1029/2008JB006173.
249	[22]	H. S. Bhat, R. Dmowska, J. R. Rice, N. Kame, Dynamic Slip Transfer from the Denali to
250		Totschunda Faults, Alaska: Testing Theory for Fault Branching, Tech. Rep. 6B (2004).
251		URL http://pubs.geoscienceworld.org/ssa/bssa/article-pdf/
252		94/6B/S202/2720488/S202_946b_04601.pdf
253	[23]	X. Ma, A. Elbanna, Dynamic rupture propagation on fault planes with explicit repre-
254		sentation of short branches, Earth and Planetary Science Letters 523 (2019) 115702.
255		doi:10.1016/j.epsl.2019.07.005.
256		URL https://linkinghub.elsevier.com/retrieve/pii/
257		S0012821X19303887
		50012021115000007
258	[24]	R. Feyiz Kartal, F. Tuba Kadirioğlu, Kinematic of East Anatolian Fault and Dead Sea
258 259	[24]	R. Feyiz Kartal, F. Tuba Kadirioğlu, Kinematic of East Anatolian Fault and Dead Sea Fault, Tech. rep. (2013).
258 259 260	[24]	R. Feyiz Kartal, F. Tuba Kadirioğlu, Kinematic of East Anatolian Fault and Dead Sea Fault, Tech. rep. (2013). URL https://www.researchgate.net/publication/271852091
258 259 260 261	[24]	 R. Feyiz Kartal, F. Tuba Kadirioğlu, Kinematic of East Anatolian Fault and Dead Sea Fault, Tech. rep. (2013). URL https://www.researchgate.net/publication/271852091 D. J. Andrews, Rupture Velocity of Plane Strain Shear Cracks., J Geophys Res 81 (32)
258 259 260 261 262	[24]	 R. Feyiz Kartal, F. Tuba Kadirioğlu, Kinematic of East Anatolian Fault and Dead Sea Fault, Tech. rep. (2013). URL https://www.researchgate.net/publication/271852091 D. J. Andrews, Rupture Velocity of Plane Strain Shear Cracks., J Geophys Res 81 (32) (1976) 5679–5687. doi:10.1029/JB081i032p05679.
258 259 260 261 262 263	[24] [25]	 R. Feyiz Kartal, F. Tuba Kadirioğlu, Kinematic of East Anatolian Fault and Dead Sea Fault, Tech. rep. (2013). URL https://www.researchgate.net/publication/271852091 D. J. Andrews, Rupture Velocity of Plane Strain Shear Cracks., J Geophys Res 81 (32) (1976) 5679–5687. doi:10.1029/JB081i032p05679. E. M. Dunham, P. Favreau, J. M. Carlson, A Supershear Transition Mechanism for Cracks,
258 259 260 261 262 263 263	[24] [25] [26]	 R. Feyiz Kartal, F. Tuba Kadirioğlu, Kinematic of East Anatolian Fault and Dead Sea Fault, Tech. rep. (2013). URL https://www.researchgate.net/publication/271852091 D. J. Andrews, Rupture Velocity of Plane Strain Shear Cracks., J Geophys Res 81 (32) (1976) 5679–5687. doi:10.1029/JB081i032p05679. E. M. Dunham, P. Favreau, J. M. Carlson, A Supershear Transition Mechanism for Cracks, Science 299 (5612) (2003) 1557–1559. doi:10.1126/science.1080650.
258 259 260 261 262 263 264 265	[24] [25] [26]	 R. Feyiz Kartal, F. Tuba Kadirioğlu, Kinematic of East Anatolian Fault and Dead Sea Fault, Tech. rep. (2013). URL https://www.researchgate.net/publication/271852091 D. J. Andrews, Rupture Velocity of Plane Strain Shear Cracks., J Geophys Res 81 (32) (1976) 5679–5687. doi:10.1029/JB081i032p05679. E. M. Dunham, P. Favreau, J. M. Carlson, A Supershear Transition Mechanism for Cracks, Science 299 (5612) (2003) 1557–1559. doi:10.1126/science.1080650. URL https://www.science.org/doi/10.1126/science.1080650
258 259 260 261 262 263 264 265 266	[24] [25] [26]	 R. Feyiz Kartal, F. Tuba Kadirioğlu, Kinematic of East Anatolian Fault and Dead Sea Fault, Tech. rep. (2013). URL https://www.researchgate.net/publication/271852091 D. J. Andrews, Rupture Velocity of Plane Strain Shear Cracks., J Geophys Res 81 (32) (1976) 5679–5687. doi:10.1029/JB081i032p05679. E. M. Dunham, P. Favreau, J. M. Carlson, A Supershear Transition Mechanism for Cracks, Science 299 (5612) (2003) 1557–1559. doi:10.1126/science.1080650. URL https://www.science.org/doi/10.1126/science.1080650 Y. Liu, N. Lapusta, Transition of mode II cracks from sub-Rayleigh to intersonic speeds in
258 259 260 261 262 263 264 265 266 266	[24] [25] [26]	 R. Feyiz Kartal, F. Tuba Kadirioğlu, Kinematic of East Anatolian Fault and Dead Sea Fault, Tech. rep. (2013). URL https://www.researchgate.net/publication/271852091 D. J. Andrews, Rupture Velocity of Plane Strain Shear Cracks., J Geophys Res 81 (32) (1976) 5679–5687. doi:10.1029/JB081i032p05679. E. M. Dunham, P. Favreau, J. M. Carlson, A Supershear Transition Mechanism for Cracks, Science 299 (5612) (2003) 1557–1559. doi:10.1126/science.1080650. URL https://www.science.org/doi/10.1126/science.1080650 Y. Liu, N. Lapusta, Transition of mode II cracks from sub-Rayleigh to intersonic speeds in the presence of favorable heterogeneity, Journal of the Mechanics and Physics of Solids

12

[28] Y. Huang, J.-P. Ampuero, Pulse-like ruptures induced by low-velocity fault zones, Journal
 of Geophysical Research 116 (B12) (2011) B12307. doi:10.1029/2011JB008684.
 URL http://doi.wiley.com/10.1029/2011JB008684

- [272 [29] X. Ma, A. Elbanna, Effect of off-fault low-velocity elastic inclusions on supershear rupture dynamics, Geophysical Journal International 203 (1) (2015) 664–677.
 doi:10.1093/gji/ggv302.
- 275 URL https://academic.oup.com/gji/article-lookup/doi/10. 276 1093/gji/ggv302
- [30] D. J. Andrews, Y. Ben-Zion, Wrinkle-like slip pulse on a fault between different materials,
 Journal of Geophysical Research B: Solid Earth 102 (B1) (1997) 553–571. doi:10.
 1029/96jb02856.
- [31] H. S. Bhat, R. L. Biegel, A. J. Rosakis, C. G. Sammis, The effect of asymmetric damage on
 dynamic shear rupture propagation II: With mismatch in bulk elasticity, Tectonophysics
 493 (3-4) (2010) 263–271. doi:10.1016/j.tecto.2010.03.016.
- [32] M. Abdelmeguid, A. Elbanna, Sequences of seismic and aseismic slip on
 bimaterial faults show dominant rupture asymmetry and potential for elevated seismic hazard, Earth and Planetary Science Letters 593 (2022) 117648.
 doi:10.1016/j.epsl.2022.117648.
- 287 URL https://linkinghub.elsevier.com/retrieve/pii/ 288 S0012821X22002849

13



Figure 1: Map of the East Anatolian Fault (EAF) zone highlighting the estimated location of the hypocenter of the $M_w 7.8$ Kahramanmaraş earthquake. The dashed line represents the inferred splay fault trace based on the recorded seismicity obtained from AFAD. The green diamonds indicate the location of the nearest seismic station to the fault trace. The black arrows indicate the left lateral sense of motion of the fault. The insert is a schematic of the relative epicenteral and hypocentral locations of the stations.



Figure 2: Supershear characteristics of near field records at stations TK:NAR, and KO:KHMN. (a) The instrument corrected records of the fault parallel, fault normal, and vertical particle velocities obtained at stations TK:NAR (black solid line), and KO:KHMN (red solid line). Note that the fault parallel component is larger than the fault normal component suggesting supershear rupture propagation. The blue dashed line indicates the arrival of the P-wave, the red dashed line indicates the arrival of the shear Mach front, and the black dashed line indicates the arrival of the trailing Rayleigh signature. (b) The theoretical relationship between the ratios of FP and FN velocity changes due the passage of the Mach front and supershear rupture speed normalized by the shear wave speed. For a ratio of velocity changes ~ 1.2 , the rupture propagates at approximately 1.55Cs, (c) Schematic diagram showing the top view on the surface highlighting the location of the stations, as well as the arrival of the shear Mach front. The green triangles indicate the locations of the stations. The epicenter is marked by a yellow star. The transition point is marked by the green square and associated error bars. The green arrow indicates the rupture propagation direction.



Figure 3: The transition from sub-Rayleigh to supershear rupture propagation is captured by the TK:4615 station. (a) The instrument corrected records of the fault parallel, fault normal, and vertical particle velocities. The highlighted region indicates the emergence of a supershear pulse ahead of the characteristic signature of a sub-Rayleigh rupture. (b) A schematic of the location of the station relative to the epicenter and hypocenter (yellow stars) location. The green triangle indicates the location of the stations. The epicenter is marked by a yellow star. The transition point is marked by the green square and associated error bars. The green arrow indicates the rupture propagation direction. Station TK:4615 is located within close proximity to the transition point.